



Mass Balance of 14 Icelandic Glaciers, 1945–2017: Spatial Variations and Links With Climate

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To date, most mass balance studies in Iceland have concentrated on the three largest ice caps. This study turns the focus toward smaller Icelandic glaciers, presenting geodetic mass-balance estimates for 14 of them (total area 1,005 km² in 2017) from 1945 to 2017, in decadal time spans. These glaciers, distributed over the country, are subject to different climatic forcing. The mass balance, derived from airborne and spaceborne stereo imagery and airborne lidar, is correlated with precipitation and air temperature by a first-order equation including a reference-surface correction term. This permits statistical modeling of annual mass balance, used to temporally homogenize the mass balance for a region-wide mass balance assessment for the periods 1945–1960, 1960–1980, 1980–1994, 1994–2004, 2004–2010, and 2010–2017. The 14 glaciers were close to equilibrium during 1960–1994, with an area-weighted mass balance of 0.07 ± 0.07 m w.e. a⁻¹. The most negative mass balance occurred in 1994–2010, accounting for -1.20 ± 0.09 m w.e. a⁻¹, or 21.4 ± 1.6 Gt (1.3 ± 0.1 Gt a⁻¹) of mass loss. Glaciers located along the south and west coasts show higher decadal mass-balance variability and static mass-balance sensitivities to summer temperature and winter precipitation, -2.21 ± 0.25 m w.e. a⁻¹ K⁻¹ and 0.22 ± 0.11 m w.e. a⁻¹(10%)⁻¹, respectively, while glaciers located inland, north and northwest, have corresponding mass-balance sensitivities of -0.72 ± 0.10 m w.e. a⁻¹ K⁻¹ and 0.13 ± 0.07 m w.e. a⁻¹(10%)⁻¹. These patterns are likely due to the proximity to warm (south and west) vs. cold (northwest) oceanic currents.

Keywords: region-wide mass balance, glacier-climate relationship, mass-balance sensitivity, Iceland, remote sensing, historical aerial photographs

1. INTRODUCTION

Glacier mass balance is a robust proxy closely linked to climate variations (Ahlmann, 1940; Ohmura, 2011; Vaughan et al., 2013; Bojinski et al., 2014). At high latitudes, mass balance is related to winter precipitation (winter snow) and summer temperature (a proxy for available energy to melt snow and ice). Glaciers have variable response times to changing climate, ranging from a few years to several decades depending on their thickness, slope and mass turnover (Jóhannesson et al., 1989; Lüthi and Bauder, 2010; Harrison, 2013; Roe et al., 2017). They filter out the high frequency seasonal and shorter-term climate variability, resulting in length or volume changes that represent

integrated response to longer-term climate variability (Elsberg et al., 2001; Marzeion et al., 2012; Christian et al., 2018).

Glaciological mass-balance observations are sparse and costly, whereas mass balance inferred from remote-sensing observations, e.g., geodetic mass balance (Cogley et al., 2011), has become the most common method to measure glacier-mass changes in glacierized regions without demanding field logistics (Zemp et al., 2019). The remote sensing era started during the early 1900s, and the mapping cameras developed rapidly in the 1930s, leading to numerous airborne and spaceborne photogrammetric and photoreconnaissance surveys worldwide (Livingston, 1963; Spriggs, 1966; Bindschadler and Vornberger, 1998). These surveys provide valuable sources to create Digital Elevation Models (DEMs) with the potential for geodetic mass-balance measurements (e.g., Finsterwalder, 1954; Bolch et al., 2011; Magnússon et al., 2016a; Fieber et al., 2018).

Despite the vast amount of historical archives of stereo images, the geodetic records of the first half of the twentieth century are still scarce and mostly based on contour maps (e.g., Bauder et al., 2007). Observations became fairly common after 1980 (e.g., Fischer et al., 2015) and have been very frequent since 2000 (e.g., Zemp et al., 2019) due to the rapid development and availability of sensors with capabilities to measure the glacier surface geometry (e.g., optical stereoscopic imagery, radar, and lidar).

Spatially distributed geodetic mass balance is available for several glacierized regions, e.g., the Alps (Fischer et al., 2015; Huss et al., 2015; Berthier et al., 2016), Andes (Soruco et al., 2009; Braun et al., 2019; Dussaillant et al., 2019), Greenland (Gardner et al., 2013; Huber et al., 2020), Norway (Andreassen et al., 2020), and High Mountain Asia (Kääb et al., 2012; Brun et al., 2017; Shean et al., 2020). These studies advance the understanding of the relationship between glacier variations and climate, are useful to calibrate regional and global climate models and constrain regional glacier-mass loss and sea-level rise (e.g., Marzeion et al., 2014; Huss et al., 2015).

In Iceland, mass-balance observations have mostly focused on the three largest ice caps: Vatnajökull, Langjökull, and Hofsjökull (7,676, 844, and 813 km² in 2017, respectively; Hannesdóttir et al., submitted), with a 30-year record of glaciological mass balance (Björnsson and Pálsson, 2008; Pálsson et al., 2012; Björnsson et al., 2013; Thorsteinsson et al., 2017; Aðalgeirsdóttir et al., submitted). These account for about 90% of the total glacierized area, and their mass-balance records have been used to estimate the Icelandic glacier-mass loss and sea-level rise contribution (Björnsson et al., 2013). Other glaciers and ice caps are less significant for the total Icelandic mass loss, but they are spatially distributed across all parts of Iceland and have the potential to provide insights into regional climate variations.

The aim of this study is to produce a catalog of maps of elevation difference and a 70-year record of geodetic mass balance of spatially distributed glaciers in Iceland. The mass balance is statistically correlated to records of temperature and precipitation to infer its static sensitivity to climate fluctuations, and the mass balance is temporally homogenized for a region-wide multi-temporal mass-balance study. The results

are discussed in the context of climate-driven changes of the 14 glaciers.

2. STUDY AREAS

For the current study, 14 glaciers and ice caps were selected, distributed across all parts of Iceland (**Figure 1**). They are located in different climatic regimes: the climate near the southern coast is influenced by the warm Irminger oceanic current, whereas the climate at the northern coast is affected by the cold East Greenland oceanic current (e.g., Björnsson and Pálsson, 2008; Björnsson et al., 2013). Geodetic mass-balance estimates have recently been obtained for three of the glaciers: Drangajökull (Magnússon et al., 2016a), Tungnafellsjökull (Gunnlaugsson, 2016), and Eyjafjallajökull (Belart et al., 2019). Glaciological mass-balance observations have been carried out at a few locations on Mýrdalsjökull since 2001 (Ágústsson et al., 2013) and on Drangajökull during 2005–2015 (e.g., Belart et al., 2017; Anderson et al., 2018). Previous studies have also calculated geodetic mass balance over one to two decades on Snæfellsjökull (Jóhannesson et al., 2011), Eyjafjallajökull, Tindfjallajökull, and Torfajökull (Guðmundsson et al., 2011). The rest of the glaciers have very limited, or no, previous mass-balance observations.

The size of the glaciers varies from 3 km² for Hofsjökull Eystri to 517 km² for Mýrdalsjökull, and their total area is 1,005 km², or 9.8% of the total of Icelandic glaciers in 2017; Hannesdóttir et al., submitted. Their elevation span ranges from ~200 m (Hofsjökull Eystri) to ~2,000 m (Öræfajökull). Further characteristics of the studied glaciers are shown in **Table 1**.

The three largest ice caps, Vatnajökull, Langjökull, and Hofsjökull, were excluded from this study with the exception of Öræfajökull, which is a part of Vatnajökull (**Figure 1**). This was due to the complexity of the required processing for those large glaciers: the relatively small footprint of the historical aerial photographs in comparison with the glacier area would limit the availability of bare ground areas needed for reference (i.e., in the vicinity of the ice cap and nunataks) for a large amounts of aerial photographs, causing large distortions in the resulting DEMs. The aerial surveys of these ice caps were also carried out over multiple dates (months to years), complicating the mosaicking and interpretation of the results.

3. DATA

The data used in this study are described by Belart et al. (2019), and consists of a large collection of stereoscopic imagery available in Iceland from 1945 to 2017, from airborne and spaceborne, frame camera and pushbroom sensors, together with airborne lidar data (**Figure 2**).

A total of 836 aerial photographs acquired during 1945–1995 were collected from the National Land Survey of Iceland. The interval between surveys for each glacier was ~10–20 years (e.g., Magnússon et al., 2016a; Pedersen et al., 2018; Belart et al., 2019). An additional series of photographs of Mýrdalsjökull and Torfajökull (acquired in 1999) were obtained from the private company Loftmyndir.

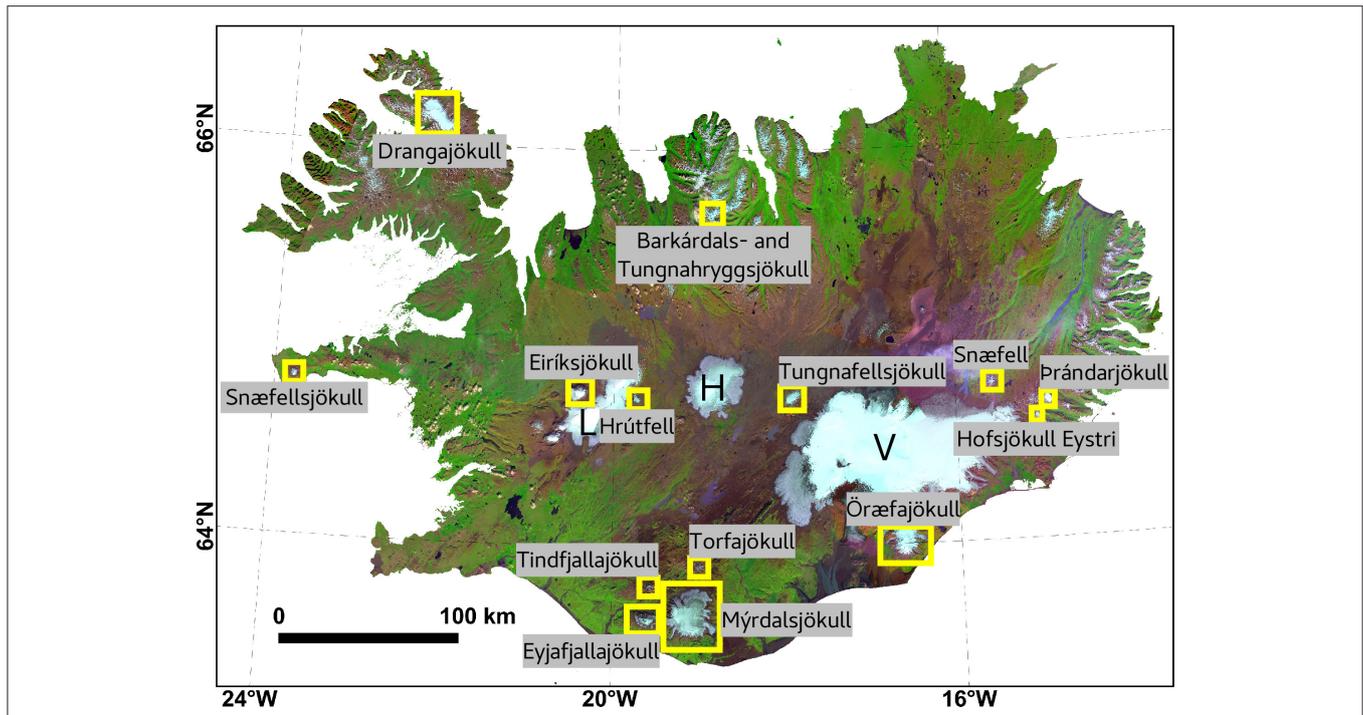


FIGURE 1 | Map of Iceland showing the 14 glaciers of this study (yellow boxes), as well as the three largest ice caps: Langjökull (L), Hofsjökull (H), and Vatnajökull (V). Background: composite mosaic from Landsat 8 (www.usgs.gov), created by the National Land Survey of Iceland (www.lmi.is).

TABLE 1 | Overview of 14 glaciers and ice caps.

	Area (km ²)	Elevation (m a.s.l.)	Mean thickness
Barkárdals- and Tungnahryggsjökull	17.7	762–1,362	N/A
Drangajökull	137.6	60–910	~100 m (Magnússon et al., 2016b)
Eiríksjökull	18.6	570–1,670	N/A
Eyjafjallajökull	65.5	200–1,630	N/A
Hofsjökull Eystri	3.0	890–1,130	N/A
Hrútfell	4.2	690–1,370	N/A
Mýrdalsjökull	517.0	170–1,490	~230 m (Björnsson et al., 2000)
Öraefajökull	163.2	0–2,110	~120 m (Magnússon et al., 2012)
Snæfell	4.3	910–1,830	N/A
Snæfellsjökull	8.3	700–1,490	~50 m (Björnsson, 2017)
Tindfjallajökull	10.8	670–1,440	N/A
Torfajökull	8.1	760–1,170	N/A
Tungnafellsjökull	32.5	900–1,520	~60 m (Gunnlaugsson, 2016)
Prándarjökull	14.0	900–1,220	N/A

The areas are referred to 2017 Hannesdóttir et al., submitted.

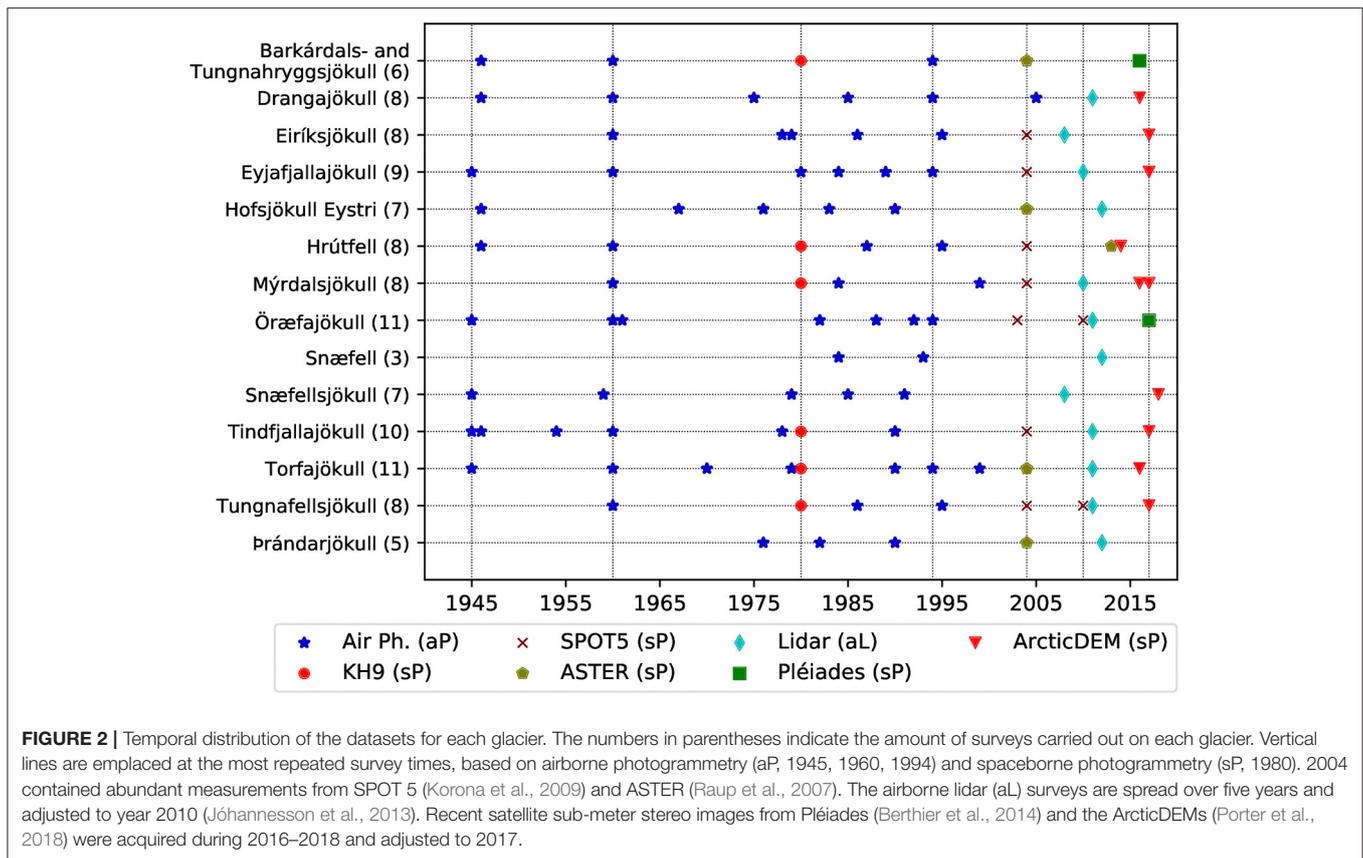
This study also used a DEM and orthoimage based on six images from Hexagon KH9 satellite acquired in August 1980, originally processed in Belart et al. (2019). They covered eight of the 14 glaciers, (from south to north) Eyjafjallajökull, Mýrdalsjökull, Tindfjallajökull, Torfajökull, Hrútfell, Tungnafellsjökull, Barkárdals-, and Tungnahryggsjökull.

In 2002–2015, SPOT 5 acquired numerous satellite stereo images of glacierized areas, particularly through the SPOT 5 Stereoscopic Survey of Polar Ice: Reference Images and Topographies (SPIRIT) project (Korona et al., 2009). This provided datapoints through the 2000s for Eyjafjallajökull, Mýrdalsjökull, Tindfjallajökull, Eiríksjökull, Hrútfell, Öraefajökull, and Tungnafellsjökull. The two last-mentioned glaciers include two acquisitions, in 2003/2004 and 2010.

Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) satellite stereo images, with modified gain setup specifically for surveying glaciers through the Global Land Ice Measurements from Space (GLIMS) project (Raup et al., 2007), were used for Barkárdals- and Tungnahryggsjökull, Torfajökull, Prándarjökull, and Hofsjökull Eystri in 2004, and for Hrútfell in 2013.

The lidar datasets were collected between 2008 and 2012, starting with Snæfellsjökull and Eiríksjökull and finishing with Snæfell and Prándarjökull (Figure 2; Jóhannesson et al., 2013). Pléiades satellite stereo images (Berthier et al., 2014) were acquired over the summers 2016 (Barkárdals- and Tungnahryggsjökull) and 2017 (Öraefajökull). Both of them were surveyed in two satellite acquisitions over the course of two weeks. Data from the ArcticDEM (Porter et al., 2018) from 2013 to 2018 were also used.

The large majority of the surveys (~80%) were carried out in August, September, or October. ~15% were carried out in July, one in June and one in November. Most of the surveys from 1945/1946 were done in late September or early October.



Daily gridded climatic records were used: linearly modeled precipitation from 1958 to 2007 (LT, 1 × 1 km; Crochet et al., 2007), numerically downscaled modeled precipitation from 1980–2019 (HARMONIE, 2.5 × 2.5 km; Nawri et al., 2017) and interpolated temperature from 1948 to 2019 based on meteorological stations (1 × 1 km; Crochet and Jóhannesson, 2011). The gridded precipitation and temperature were measured in millimeters and Celsius degrees, respectively. Remarks on the data quality and associated uncertainties of the climatic records can be found in the mentioned studies and in Belart et al. (2019).

The glacier margins of Icelandic glaciers in ~2000 were extracted from the GLIMS inventory (Raup et al., 2007), as a first approximation of the glacier outlines of the 14 glaciers.

4. METHODS

For each of the glaciers, the processing steps described by Belart et al. (2019) were followed, and we refer the readers to that article for further technical details. In summary, these consist of: (1) processing of stereo imagery and DEMs for creation of co-registered orthoimages, DEMs and difference of DEMs (dDEM); (2) glacier outline delineating for the two analyzed dates; (3) bias correction and uncertainty estimation in the volume changes using Sequential Gaussian Simulation (SGSim; Magnússon et al., 2016a); (4) filtering of outliers, gap filling and calculation of volume changes; (5) applying a seasonal

correction to account for the volume change between the date of each survey and 1 October of the respective year; (6) geodetic mass-balance estimation, using a conversion factor of $850 \pm 60 \text{ kg m}^{-3}$ between volume and mass change (Huss, 2013). This workflow resulted in two values of geodetic mass balance, \hat{B} , one relative to the survey dates (floating-date system) and one fixed to the hydrological year, 1 October to 30 September (fixed-date system).

The photogrammetric processing of the aerial photographs (step 1) was done with MicMac software (Pierrot Deseilligny and Clery, 2011; Rupnik et al., 2017) using Ground Control Points (GCPs) extracted from a reference, in most cases from a lidar DEM, except for Barkárdals- and Tungnahryggsjökull (Pléiades DEM and orthoimage, Papasodoro et al., 2015), and Hróútfell (ArcticDEM). The pushbroom stereo images, i.e., Pléiades, ASTER, and SPOT 5 data, were processed using Ames StereoPipeline (ASP) software (Shean et al., 2016, 2020). The resulting DEMs and orthoimages, as well as the ArcticDEMs, were co-registered to the aforementioned reference using the Iterative Closest Point method in ASP (Shean et al., 2016).

The glacier outlines obtained from GLIMS were updated using the orthoimages (shaded relief images of the DEMs when no orthoimages were available). We modified the definition of the margins of N-Eiríksjökull, where we included the debris-covered ice in the low areas (Figure S3). When gaps in the orthoimages were present at particular locations of the margins, due to clouds or bad stereo coverage, the outline was completed using aerial

photographs, Landsat, or ASTER data with dates differing up to 2 years from the year-of-interest.

The SGSim method (Magnússon et al., 2016a), used for retrieving uncertainties and bias correction in the volume changes, required as input the ice-free areas of the dDEMs, manually masked from outliers. Semivariograms were created from the dDEMs on ice-free areas and manually fitted into semivariogram models. A total of 1,000 SGSims were run on each glacier, obtaining a map of errors propagated onto the respective glacier. These maps of errors were then subtracted from the dDEMs to locally remove systematic errors in elevation, and a histogram retrieved from the SGSim was used to extract the uncertainty of the volume change obtained from the dDEMs (95% confidence level).

The seasonal correction applied is based on a degree-day melt model combined with a late-summer snow accumulation model, using the daily gridded climate records. The setup parameters and uncertainties of the model are described in Belart et al. (2019). This seasonal correction model was run using bootstrapping, and from the 1,000 realizations of the runs we obtained a volume change from the date of survey to the 1 Oct, as well as the associated uncertainty (95% confidence level).

In some cases, we combined DEMs from surveys carried out up to two years apart, in order to fill gaps in the maps of elevation difference. This was done for Tindfjallajökull (70% in Sep 1945 and 30% in Sep 1946), Örafajökull (60% in Aug 1960 and 40% in Jul 1960, and 50% in Aug 1992 and 50% in Aug 1994), Eiríksjökull (80% in Aug 1978 and 20% in Aug 1979), and Mýrdalsjökull (90% in 2016 and 10% in 2017). Each of the surveys covered the majority of the elevation span of the respective ice cap. In these cases the dDEM with smaller coverage was shifted to the main dDEM by the mean difference of the overlapping areas of the mosaic.

Remaining gaps in the maps of elevation difference (due to incomplete surveys, cloud presence, and lack of texture in the images) were interpolated as a function of elevation, using the local hypsometric method (e.g., Brun et al., 2017; McNabb et al., 2019). The uncertainty of these areas was increased based on the number of datapoints and variations in elevation difference for each elevation band (Supplement S1). The uncertainty of each geodetic mass balance was calculated as the square root of the sum of squares of the derivative of the mass balance with respect to each of the variables involved (Belart et al., 2019). The resulting uncertainty was subsequently combined with the uncertainty due to data gaps as the root sum of their squares.

Annual mass balance and static mass-balance sensitivities were calculated by correlating the fixed-date geodetic mass balance to the climatic records, using the following equation (Belart et al., 2019)

$$\dot{B} = \phi T_s + \omega P_w + \gamma \Delta \bar{A} + k, \quad (1)$$

where T_s and P_w are mean summer temperature and winter precipitation at the equilibrium line altitude (ELA), respectively. The summer is defined from 21 May to 30 September, the approximate period when the temperature is typically positive at

the ELA and the aimed dates for the glaciological mass balance campaigns in Iceland (e.g., Belart et al., 2017, 2019). The winter precipitation is normalized, i.e., divided by the average winter precipitation of the available records (Oerlemans and Reichert, 2000). The ELA is assumed to be the snowline in ~ 2004 , and unchanged over the period of study for simplicity. The term $\Delta \bar{A}$ is the difference between a reference area, e.g., 1960, and the average area between each two years of survey. Together with γ and k , this serves as a correction to relate conventional and reference-surface mass balance (Elsberg et al., 2001; Harrison et al., 2001), and gives a better correlation of mass balance with climate (Cogley et al., 2011; Huss et al., 2012). The reference-surface term is simplified by assuming that relative changes in glacier area and ice volume are linearly related for the analyzed time period (Pálsson et al., 2012; Belart et al., 2019).

Equation (1) was solved with a least-squares fit, weighted with mass-balance uncertainties (Belart et al., 2019). The coefficients ϕ and ω are equivalent to the static mass-balance sensitivity of mass balance to summer temperature ($\delta \dot{B} / \delta T_s$, m w.e. $a^{-1} K^{-1}$) and winter precipitation ($\delta \dot{B} / \delta P_w$, m w.e. $a^{-1} (10\%)^{-1}$), respectively. We solved the least-squares fit for individual glaciers, obtaining a set of ϕ , ω , γ and k for each glacier. This, however, led to unrealistically high and low sensitivity to summer temperature for Snæfellsjökull and Þrándarjökull, respectively, likely due to a limited number of observations available for these glaciers. We then grouped the glaciers in two sets, using a threshold of 1.4 m w.e. $a^{-1} K^{-1}$ from the initial sensitivities. This criteria fitted with the regional characteristics of the grouped glaciers: glaciers located at the southern and western coasts (Eyjafjallajökull, Mýrdalsjökull, Örafajökull, Snæfellsjökull, Tindfjallajökull, and Torfajökull) and glaciers in central, northern, and eastern Iceland (Drangajökull, Eiríksjökull, Hrutífell, Tungnafellsjökull, Barkárdals- and Tungnahryggsjökull, Hofsjökull Eystri, and Þrándarjökull).

We performed the least-squares fit for each group of glaciers, calculating one single coefficient ϕ and ω per group, while the coefficients γ and k were determined for each individual glacier within the group. The least-squares fit did not include Snæfell, for which only two geodetic mass balances are available (Figure 2).

For a decadal, region-wide comparison, we selected the years 1945, 1960, 1980, 1994, 2004, 2010, and 2017 as the most common survey years (Figure 2). The geodetic mass balance of each glacier was temporally homogenized (e.g., Lambrecht and Kuhn, 2007; Fischer et al., 2015) to these years when needed, by using the modeled annual mass balance from Equation (1). An uncertainty of ± 0.5 m w.e. a^{-1} was assumed for each year of modeled annual mass balance when calculating the temporally-homogenized mass balance. The annual correction was not applied to datasets acquired in 1946 due to lack of climate data.

We also estimated the geodetic mass balance for two additional, longer, time periods: 1960–1994 (relatively cold period and close to zero balance) and 1994–2010 (warmer period and the largest ice wastage). These additional geodetic mass balances were estimated using the DEMs closest in time to 1960, 1994, and 2010. They were corrected to the hydrological year, and temporally homogenized when needed, using the modeled annual mass balances.

5. RESULTS

Over 100 DEMs of the 14 glaciers, between 1945 and 2018, were utilized in this study. This included the creation of 63 DEMs from historical aerial photographs. The gaps on the glaciers were on average 15% of the total glacier area; seven DEMs contained gaps that were >30% of the glacier area, with a maximum of 40% on Snæfellsjökull in 1945 and 1959. A time series of elevation changes for each of the glaciers is shown in **Supplement S2**.

A total of 96 geodetic mass balances were computed, a number increasing to 111 after including the results from Magnússon et al. (2016a) and Belart et al. (2019), i.e., eight time periods on average for each of the 14 glaciers, and nine of them starting in 1945 (**Figure 3**). Tabulated observations of floating-date and fixed-date geodetic mass balance are provided in **Supplement S3**. These observations are also available at the World Glacier Monitoring Service (WGMS, www.wgms.ch).

Uncertainties in the geodetic estimates were typically <0.1 m w.e. a⁻¹ (95% confidence level) for periods longer than 10 years, but they increased for the shorter time periods (e.g., Huss, 2013) or if the DEMs contained significant gaps. The maximum and minimum fixed-date geodetic mass balances were, respectively, 0.84 ± 0.21 m w.e. a⁻¹ on Snæfellsjökull in 1985–1991 and -2.40 ± 0.25 m w.e. a⁻¹ on Mýrdalsjökull in 1999–2004. The seasonal correction was generally small (<0.1 m w.e. a⁻¹) over ~10-year periods, but was significantly larger for the shorter time periods (e.g., Belart et al., 2019).

The obtained mass-balance sensitivity to summer temperature was -2.21 ± 0.25 m w.e. a⁻¹ K⁻¹ for the group of high-sensitivity glaciers, and -0.72 ± 0.10 m w.e. a⁻¹ K⁻¹ for the group of low-sensitivity glaciers. The sensitivities to winter precipitation are 0.22 ± 0.11 m w.e. a⁻¹ (10%)⁻¹ for the high-sensitivity glaciers, and 0.13 ± 0.07 m w.e. a⁻¹ (10%)⁻¹ for the low-sensitivity glaciers.

The fit between mass balance and climatic variables (Equation 1) yielded a high correlation between observed and statistically derived mass balance, with a coefficient of determination $R^2 = 0.71$ for the group of high-sensitivity glaciers and $R^2 = 0.77$ for the group of low-sensitivity glaciers (**Table 2**). This correlation was substantially lower in a combined least-squares fit using all glaciers in one single group ($R^2 = 0.59$) or assuming $\gamma = 0$, i.e., without adding the reference-surface correction term (**Table 2**).

The mean (standard deviation in parenthesis) mass balance, calculated without area-weights, of the 14 glaciers was -0.42 (0.17) m w.e. a⁻¹ in 1945–1960, 0.00 (0.21) m w.e. a⁻¹ in 1960–1980, 0.12 (0.21) m w.e. a⁻¹ in 1980–1994, -1.01 (0.45) m w.e. a⁻¹ in 1994–2004, -1.22 (0.63) m w.e. a⁻¹ in 2004–2010, and -0.18 (0.22) m w.e. a⁻¹ in 2010–2017. The area-weighted mean (standard deviation in parenthesis) was 0.04 (0.12) m w.e. a⁻¹ in 1960–1980, 0.10 (0.14) m w.e. a⁻¹ in 1980–1994, -1.27 (0.53) m w.e. a⁻¹ in 1994–2004, -1.32 (0.51) m w.e. a⁻¹ in 2004–2010, and -0.18 (0.22) m w.e. a⁻¹ in 2010–2017. No area-weighted value was calculated for 1945–1960 due to the lack of observations for Mýrdalsjökull (by far the largest glacier among the 14) in this time period. Tabulated results of the temporally homogenized geodetic mass balance and associated uncertainties (95% confidence level) are shown in **Supplement S4**.

The overall results from 1960 to 1994 indicate a near-zero balance for the 14 glaciers. Most of the mass loss occurred in 1994–2010: 21.4 ± 1.6 Gt (1.3 ± 0.1 Gt a⁻¹), or 0.059 ± 0.005 mm Sea Level Equivalent (SLE). In this time period, Mýrdalsjökull accounted for about 70% of the mass loss. The group of glaciers <100 km² contributed significantly more to the mass loss than some larger (>100 km²) ice caps such as Örfajökull, and more than twice as much as Drangajökull for the same period (**Table 3**).

6. DISCUSSION

6.1. Spatio-Temporal Mass Balance and Climate Distribution

The long (>100-year) temperature record from Stykkishólmur (W-Iceland) indicates a maximum in the 1930s followed by gradual cooling until the 1960s, and then warming again in the 1990s. During 1945–1960, the glaciers experienced mass losses (**Figure 4A**). During 1960–1980 most of the glaciers had a close to zero mass balance, with the exception of Torfajökull and Mýrdalsjökull, which had slightly negative mass balance. Most glaciers gained mass or were close to equilibrium in 1980–1994, with the highest mass gain on Snæfellsjökull (W) and Eyjafjallajökull (S) (**Figures 3, 4B,C, 5**). These observations of mass balance during 1945–1990 agree with previous studies on several Icelandic glaciers (Pálsson et al., 2012; Björnsson et al., 2013; Hannesdóttir et al., 2015).

A substantial mass loss was experienced by the 14 glaciers in 1994–2010 (**Figures 4D,E**). The highest mass loss is found at coastal glaciers in the south and west and glaciers located at lower elevations (lower than ca. 1,200 m a.s.l., **Table 1**). The area-weighted mass balance of the 14 glaciers was -1.27 ± 0.09 m w.e. a⁻¹ in this period (**Table 3**), similar to the glaciological mass balance measured at Hofsjökull and Langjökull (-1.3 m w.e. a⁻¹), but more negative than for Vatnajökull (-0.8 m w.e. a⁻¹) (Pálsson et al., 2012; Björnsson et al., 2013; Thorsteinnsson et al., 2017) during the same period. The presented results for the periods 1994–2010 and 2004–2010 also encompass the effects of the increased summer melt of 2010 as a consequence of the Eyjafjallajökull eruption in 2010 (e.g., Björnsson et al., 2013).

There was less mass loss in 2010–2017 than in the previous two decades (**Figure 4F**), as has been also observed in Greenland (Shepherd et al., 2020), in mainland Norway (Andreassen et al., 2020), and also in Iceland by gravimetric methods (Wouters et al., 2019) and by glaciological observations (Aðalgeirsdóttir et al., submitted). Örfajökull gained mass during this period (0.43 ± 0.08 m w.e. a⁻¹) and some of its outlets have shown significant advances (**Figure S9**). It covers the largest elevation range (0–2,100 m a.s.l.), collects the highest amount of precipitation in Iceland (Crochet et al., 2007) with likely summer snowfalls and has steep outlets leading to rapid mass transport. The modeled precipitation used in our study (Nawri et al., 2017) indicates increased winter precipitation in recent years on Örfajökull.

The decadal variability of mass balance is strongly related to the proximity of the glaciers to the coasts (**Figure 6**). Glaciers located at the south and west coast are classified as maritime (e.g., De Woul and Hock, 2005) and show large mass-balance

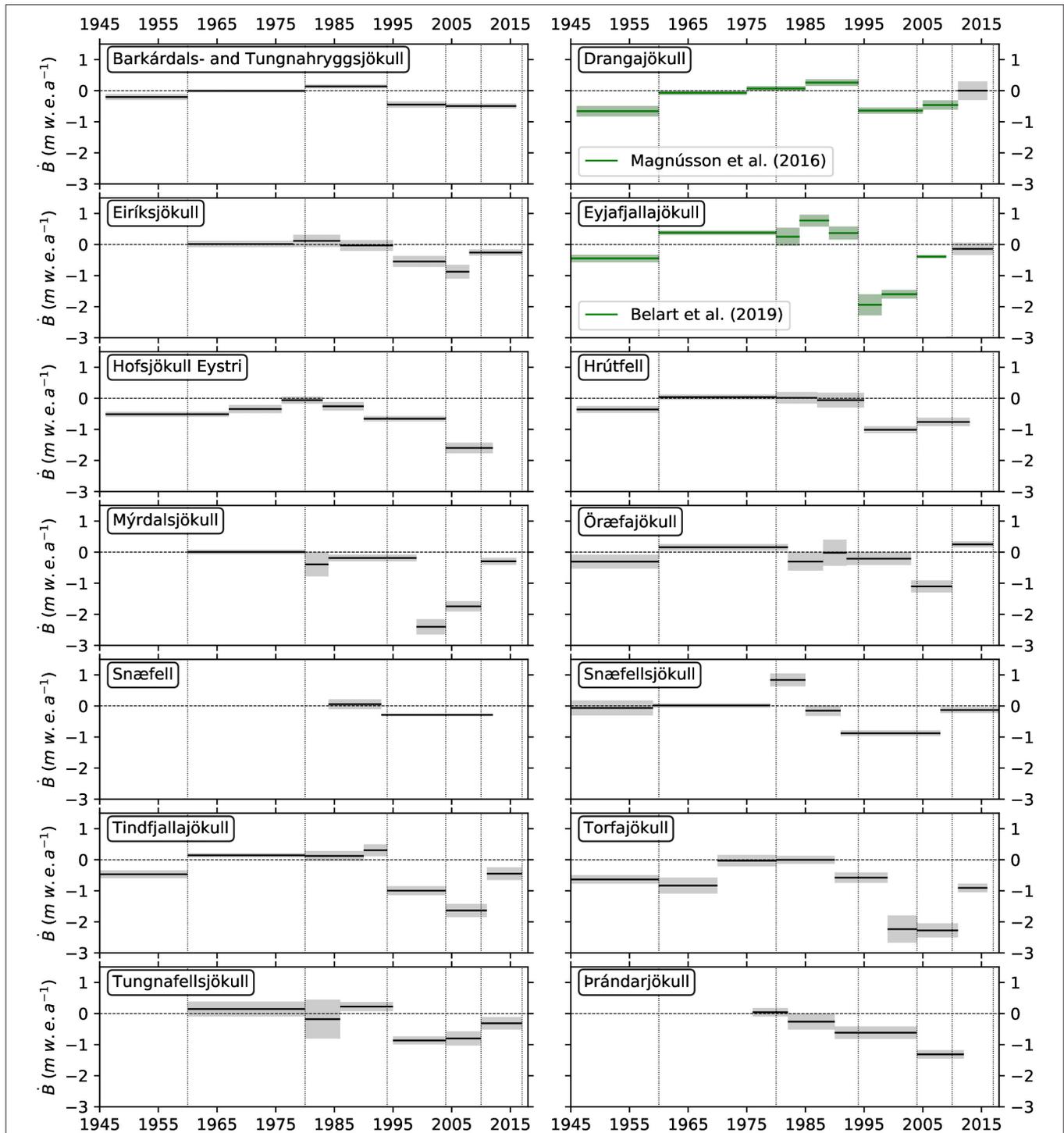


FIGURE 3 | Glacier-wide geodetic mass balance of the 14 glaciers. The mass balance is estimated from 1 October to 30 September (fixed-date system), based on seasonal corrections. Gray bands indicate uncertainty at 95% confidence level (Magnússon et al., 2016a). Dashed vertical lines show the years selected for a region-wide, temporally homogenized mass-balance comparison. Additional mass balances can be found in **Table S1**, relative to longer time periods and for Torfajökull (1970–1979 and 1979–1990) and Tindfjallajökull (1960–1978 and 1978–1980), which are further analyzed in the discussion.

variations and high static sensitivities to summer temperature and winter precipitation. The inland glaciers are subject to rain shadows and experience less variable precipitation. Their

temporal variability in mass balance may be explained by differences in elevation. Torfajökull and Tindfjallajökull have a narrower elevation span, larger decadal mass-balance variability

TABLE 2 | Mass-balance sensitivities of the glaciers to a 1°C rise in temperature and to a 10% increase in precipitation.

	$\delta\dot{B}/\delta T_s$ (m w.e. a ⁻¹ K ⁻¹)	$\delta\dot{B}/\delta P_w$ (m w.e. a ⁻¹ (10%) ⁻¹)	R ²	R ² ($\gamma = 0$)
All glaciers	-1.10 ± 0.14	0.17 ± 0.08	0.59	0.53
High-sensitivity glaciers (South and West)	-2.21 ± 0.25	0.22 ± 0.11	0.71	0.67
Low-sensitivity glaciers (North, inland and East)	-0.72 ± 0.10	0.13 ± 0.07	0.79	0.57

Uncertainties (1σ) were extracted from the variance matrix as result of the least-squares fit. The high-sensitivity glaciers comprise Eyjafjallajökull, Mýrdalsjökull, Öraefajökull, Snæfellsjökull, Tindfjallajökull, and Torfajökull. The low-sensitivity glaciers comprise Barkárdals- and Tungnahryggjökull, Drangajökull, Eiríksjökull, Hofsjökull Eystri, Hróttfell, Tungnafellsjökull, and Þrándarjökull.

TABLE 3 | Mass changes of 14 Icelandic glaciers from 1945–1960 to 1994–2010.

	Area (2017) (km ²)	\dot{B}_{1945}^{1960} (m w.e. a ⁻¹) [Gt]*	\dot{B}_{1960}^{1994} (m w.e. a ⁻¹) [Gt]	\dot{B}_{1994}^{2010} (m w.e. a ⁻¹) [Gt]
Drangajökull	137.6	-0.66 ± 0.17 [-1.5 ± 0.4]	0.13 ± 0.04 [0.7 ± 0.2]	-0.62 ± 0.09 [-1.4 ± 0.2]
Mýrdalsjökull	517.0	N/A [N/A]	0.02 ± 0.05 [0.4 ± 1.0]	-1.48 ± 0.17 [-13.9 ± 1.6]
Öraefajökull	163.2	-0.31 ± 0.23 [-0.8 ± 0.6]	0.09 ± 0.07 [0.6 ± 0.5]	-1.01 ± 0.1 [-2.8 ± 0.3]
Others**	187.0	-0.42 ± 0.07 [-1.0 ± 0.2]	0.16 ± 0.02 [1.3 ± 0.1]	-0.99 ± 0.05 [-3.4 ± 0.2]
Total	1,004.8	N/A [N/A]	0.07 ± 0.03 [2.9 ± 1.1]	-1.20 ± 0.09 [-21.4 ± 1.6]

The table shows area (km², Hannesdóttir et al., submitted), specific (m w.e. a⁻¹) and total (Gt, in brackets) mass change rates for the three largest glaciers and for the sum of the other glaciers (glaciers <100 km²) as investigated in this study. The values correspond to the fixed-date geodetic mass balance, after temporal homogenization when needed. Mass balance and mass loss involving multiple glaciers is calculated using area-weights. *Values used for Mýrdalsjökull, Eiríksjökull, and Tungnafellsjökull reach only back to 1960. **Þrándarjökull and Snæfell are excluded in this analysis, due to limited data coverage.

and are more sensitive to temperature changes than other inland glaciers, hence they were included in the coastal glaciers group. Glaciers farther north, like Drangajökull and the glaciers on Tröllaskagi, have significantly lower mass-balance decadal variability and sensitivities and are situated close to the cold East Greenland oceanic current (Björnsson et al., 2013). Þrándarjökull and Hofsjökull Eystri, despite being located near the SE coast, show lower sensitivity to summer temperature and were grouped with the central and northern glaciers.

The regional pattern of high decadal mass-balance variability and sensitivities also fits the regional pattern of average melt-season albedo during 2000–2019 estimated for most Icelandic glaciers by Gunnarsson et al. (2020). The southern glaciers have significantly lower average melt-season albedo than the glaciers located in central or northern Iceland. This can be partly explained by the two aforementioned climatic regimes: higher summer temperatures in the southern glaciers lead to stronger

melt and an early appearance of the firn and ice layer, and the opposite effect would occur in the central and northern glaciers. However, this is also partly explained by the presence of more dust and tephra on the southern glaciers, which can also significantly lower the albedo (e.g., Möller et al., 2014, 2019; Gunnarsson et al., 2020).

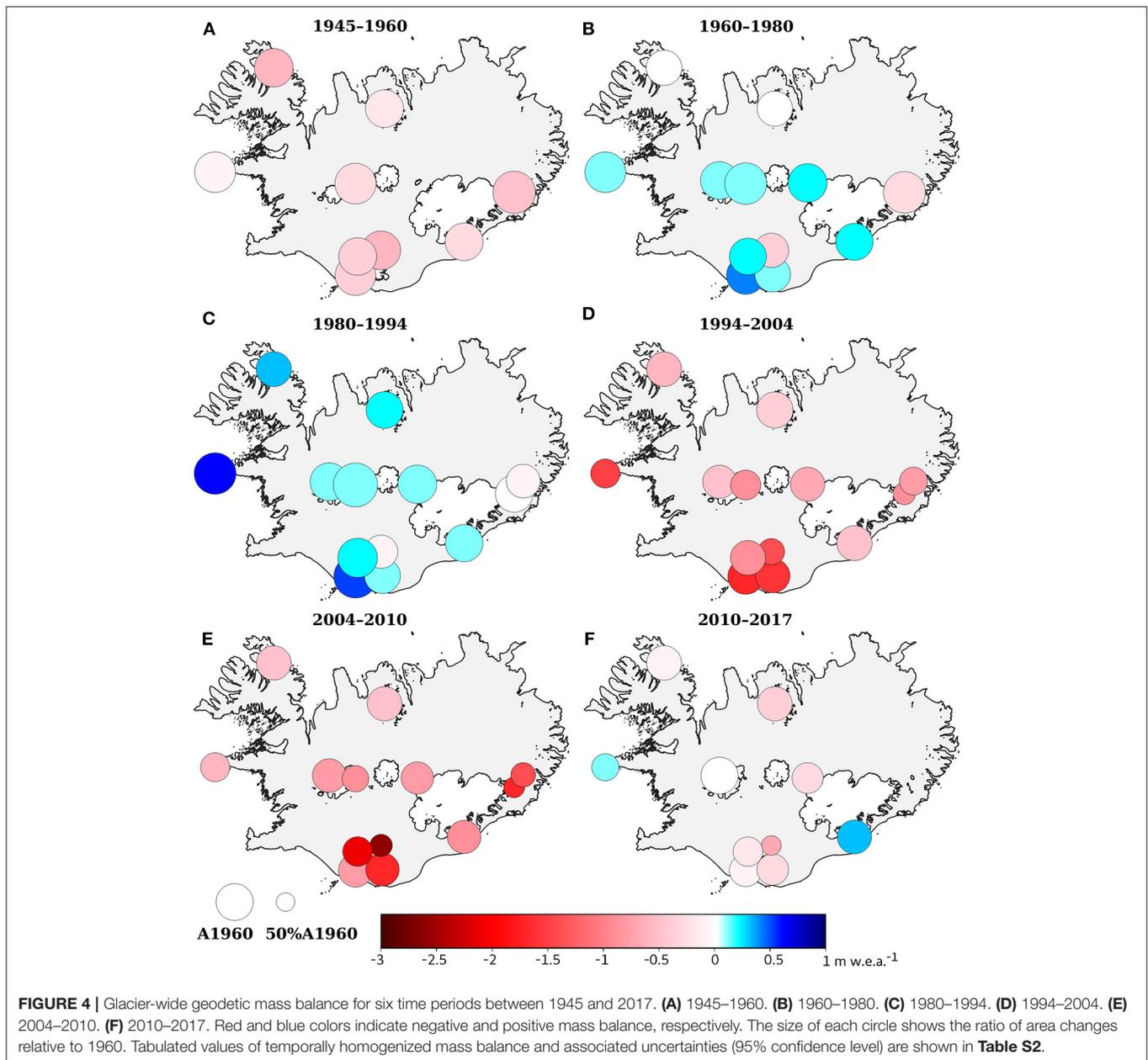
We observed high intraregional variability of mass balance, particularly prominent for Tindfjallajökull, Torfajökull, Eyjafjallajökull, and Mýrdalsjökull (S-Iceland, within 40 km of each other). Analogously, different catchments of some of the glaciers can exhibit substantial differences in mass balance: (1) Drangajökull shows strong differences in mass balance between the eastern and western catchments, probably associated with the effects of precipitation and snow drift (Magnússon et al., 2016a,b; Belart et al., 2017). (2) On Mýrdalsjökull, the northern (inland) outlets show an almost continuous lowering during all observed time periods, as opposed to the southern (coastal) outlets, which show fluctuations between lowering and thickening through the past decades (Figure S8). The same pattern is observed on Vatnajökull (Björnsson et al., 2013) and can be explained by the maritime regime of the southern outlets, with most precipitation falling at these locations, while the northern outlets are located in a rain shadow (Crochet et al., 2007; Ágústsson et al., 2013). (3) One of the outlets of Öraefajökull is affected by significant calving at its terminus (Magnússon et al., 2012), causing more negative mass balance than the rest of the ice cap (Figure S9).

Mýrdalsjökull contributed 70% of the mass loss of the studied glaciers in 1994–2010, although it accounts for only about 50% of the area encompassed by the 14 glaciers (Table 3). This shows that extrapolation of mass balance from a few glaciers to an entire region (e.g., Björnsson et al., 2013) can lead to erroneous estimates of mass loss from small glaciers. In this context, Mýrdalsjökull and the southern small glaciers contribute more to sea-level rise than the northern glaciers, like Drangajökull or the cluster of glaciers in Tröllaskagi. Nevertheless, the small glaciers cover 10% of the total glacierized area of Iceland, and their mass loss is close to one order of magnitude smaller than for the entire country (9.5 Gt a⁻¹, Björnsson et al., 2013, vs. 1.3 Gt a⁻¹).

In comparison with to long-term (>50-year) mass-balance observations in other glacierized regions, the evolution of Icelandic glaciers has followed similar patterns during the study period to those observed in the Alps (Huss et al., 2010), Pyrenees (Marti et al., 2015), or in tropical glaciers as in Cordillera Real (Sorucu et al., 2009). The intraregional variability observed in Iceland during 1994–2004 (SD = 0.44 m w.e. a⁻¹, N = 13) and 2004–2010 (SD = 0.57 m w.e. a⁻¹, N = 13) appears to be similar to that in the Himalayas in 2000–2016 (Brun et al., 2017). Some other glacierized regions such as the Alps have, however, experienced more homogeneous mass loss during the last few decades (Fischer et al., 2015).

6.2. Statistical Estimation of Mass Balance From Precipitation and Temperature Records

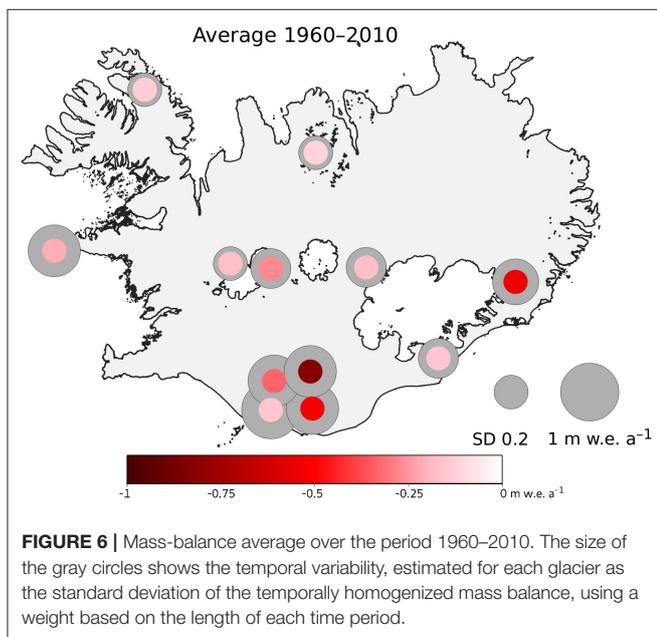
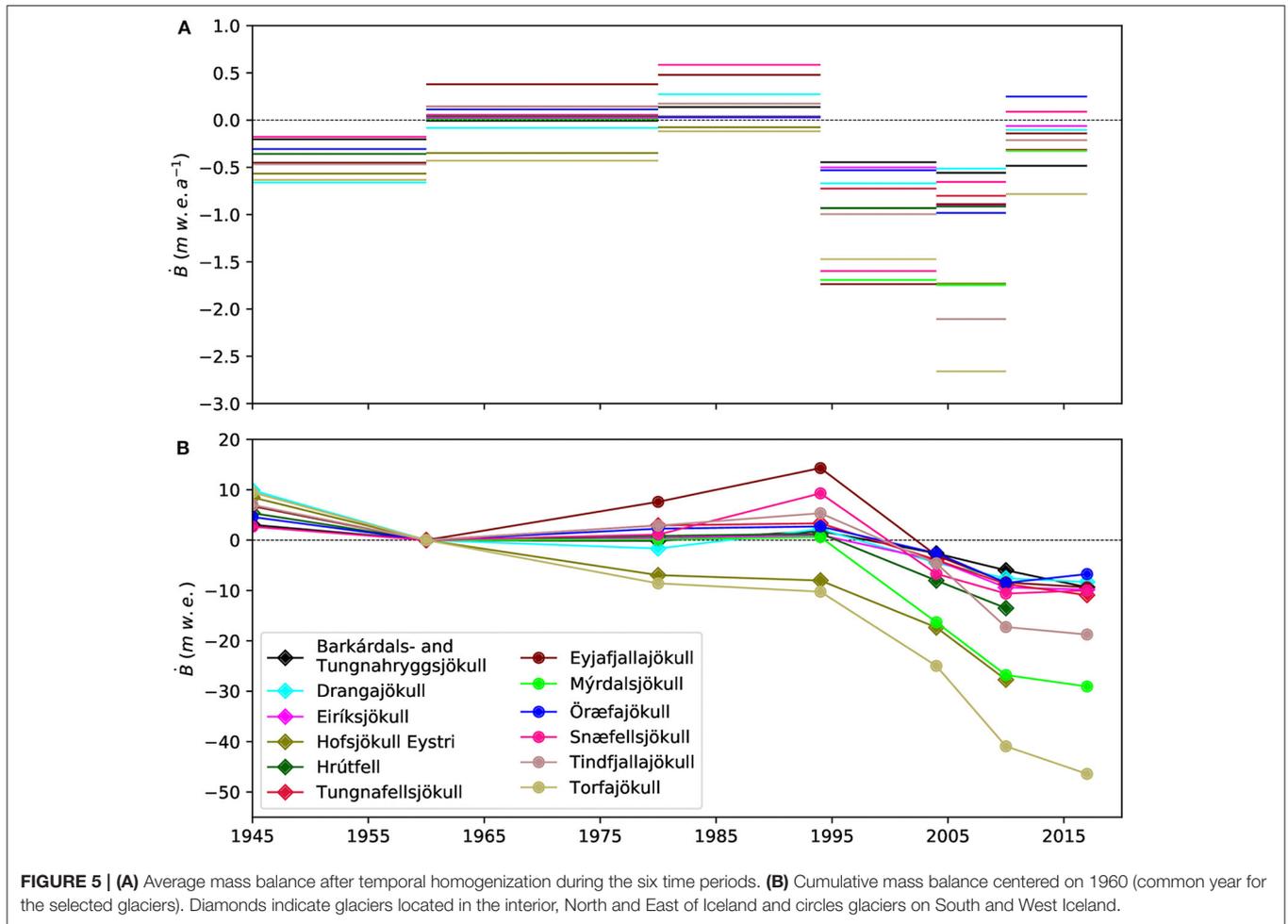
A first-order equation (Equation 1) can be used to estimate the annual mass balance as a function of summer temperature,



winter precipitation and area, permitting a practical temporal homogenization of geodetic mass balance (e.g., Lambrecht and Kuhn, 2007; Fischer et al., 2015). The mass balance correlates well with climatic variables (Cogley et al., 2011; Huss et al., 2012). In other words, a large fraction of the mass-balance variability can be explained by variations in summer temperature and winter precipitation, in addition to feedback effects due to changes in the geometry of the glacier. The reference-surface correction term improves the fit (R^2 , **Table 2**) by up to 20% for the group of central and northern glaciers.

The initial results when the Equation (1) was solved for individual glaciers (as opposed to groups of glaciers) led to unrealistically high and low mass-balance sensitivities for

Snæfellsjökull and Þrándarjökull, respectively. This can be attributed to the limited observations used for these two glaciers, but also by other factors mentioned in the following discussion. Grouping the glaciers in two regions added robustness to the statistical analysis. The obtained sensitivities agree with the range of sensitivities calculated by De Woul and Hock (2005) for specific outlet glaciers of Vatnajökull [ranging -1.5 to -2 m w.e. $a^{-1} K^{-1}$ and 0.2 to 0.3 m w.e. $a^{-1} (10\%)^{-1}$], similar to our estimates for the southern glaciers, and for specific outlet glaciers of Hofsjökull [ranging -0.7 to -1.3 m w.e. $a^{-1} K^{-1}$ and 0.1 m w.e. $a^{-1} (10\%)^{-1}$], similar to our estimates for the central and northern glaciers. Based on the joint analysis of mass balance and climate, we find that, for the



group of southern and western glaciers, the mass loss during 1994–2010 was entirely attributed to an increase in summer temperature relative to 1960–1994, while the winter precipitation remained approximately constant on average during these two time periods.

Torfajökull, with DEMs acquired in 1979 and 1980, and Tindfjallajökull, with DEMs acquired in 1978 and 1980, served as a test of the temporal homogenization. For these two glaciers, applying temporal homogenization using the 1978 and 1979 DEMs (i.e., 1 and 2 years apart from 1980) had similar results to using the 1980 DEM, with differences lower than 0.1 m w.e. a⁻¹ between observed and temporally homogenized mass balances, which is within the uncertainty of the geodetic estimates. Additionally, area-weighted averages of the annual mass balance obtained from Equation (1) were computed for the two sub-regions. These results are presented in (Figure 7), and compared with annual glaciological mass-balance observations for Vatnajökull, Langjökull and Hofsjökull (Thorsteinsson et al., 2017; WGMS, 2019; Hannesdóttir et al., submitted). The annual mass balance obtained for group of southern and western glaciers shows good correlation with the glaciological mass balance of

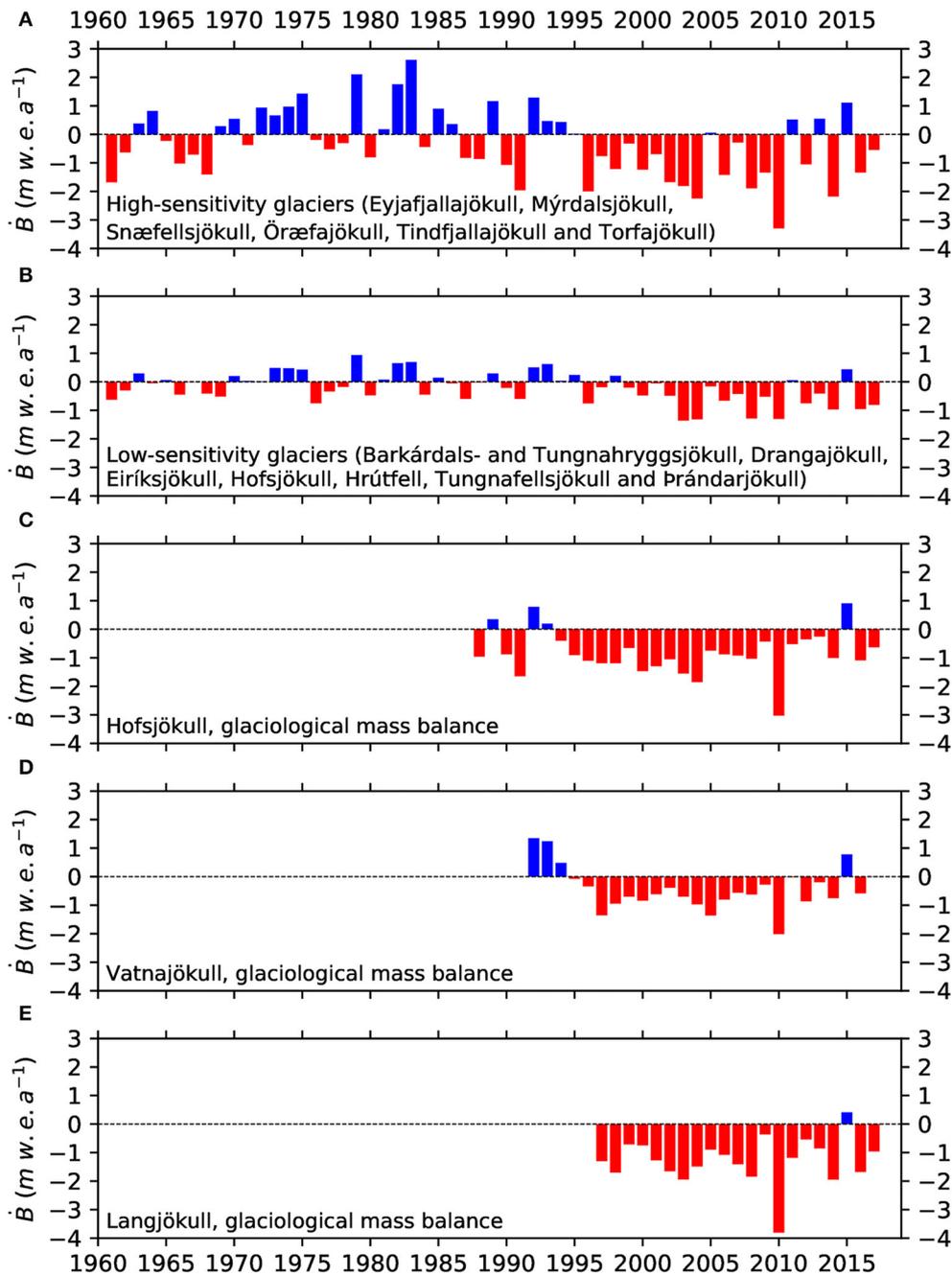


FIGURE 7 | (A,B) Area-weighted average annual mass balance of the two sub-regions, obtained from Equation (1) over the period 1960–2017. **(C–E)** Glaciological mass balance measured on Hofsjökull, Vatnajökull, and Langjökull (Thorsteinsson et al., 2017; WGMS, 2019; Aðalgeirsdóttir et al., submitted).

Vatnajökull ($R^2 = 0.52$), Langjökull ($R^2 = 0.57$), and especially with Hofsjökull ($R^2 = 0.73$).

The annual mass balance should, however, be interpreted with caution, particularly for individual glaciers and for time periods where there is a mismatch between modeled and observed mass balance (Supplement S5). The mismatches can be attributed to either an incomplete climate model or to an over-simplified linear fit between climate and mass balance. Measuring and modeling

winter precipitation is challenging (e.g., Jarosch et al., 2012). In Öræfajökull, the precipitation is overestimated by the model used (Schmidt et al., 2017). Moreover, the mass balance is controlled by other variables neglected in our simple model. Full energy-balance models can better reproduce the response of glaciers to climate variations (e.g., Arnold et al., 1996; Hock and Holmgren, 2005), accounting, for example, for albedo changes, which can intensify melting if dust (e.g., Arnalds et al., 2016; Wittmann

et al., 2017) or thin tephra (e.g., Möller et al., 2014; Gascoïn et al., 2017) is deposited on the glacier, or can decrease the melt if snow fall occurs during summer.

The mass balance of Icelandic glaciers has additional forcing besides temperature and precipitation, such as: (1) Wind-drifted snow can contribute to winter accumulation in individual catchments, as observed for Drangajökull (Magnússon et al., 2016a,b; Belart et al., 2017). (2) Debris (either rock or tephra), present on Icelandic glaciers like Eiríksjökull, enhances or reduces the melt rate depending on the thickness (e.g., Östrem, 1959). (3) Volcanic forcing can affect the mass balance, typically for one hydrological year (Björnsson et al., 2013), and possibly over longer periods (Möller et al., 2019). (4) Local changes in geothermal heat flux underneath the glacier, as observed on Mýrdalsjökull (Guðmundsson et al., 2007), can also affect the mass balance (Björnsson, 2003).

The reference-surface correction term in the mass-balance model (Equation 1) was generalized by Elsberg et al. (2001) and Harrison et al. (2001), and further in Belart et al. (2019) by assuming a linear relationship between the volume changes and area changes of the glacier. This correlation is generally found applicable in the 14 glaciers, with a typical fit $R^2 > 0.8$ between volume changes and area changes (e.g., Pálsson et al., 2012; Belart et al., 2019), but it does not apply for three surge-type catchments in Drangajökull and two in Mýrdalsjökull (Björnsson et al., 2003). Several outlet glaciers of Mýrdalsjökull and Öraefajökull, not identified as surge-type glaciers, present a series of noteworthy elevation changes during the 1960s to 1990s: their accumulation area lowered while the ablation area was gained elevation and their margins were advanced, which suggests acceleration in their ice motion. This led to a rapid area increase, while the overall volume remained unchanged (Figures S8, S9).

The suggested acceleration in the ice motion during the colder time periods, particularly as observed on Mýrdalsjökull and Öraefajökull is worth further study. This is also a key for fully describing the mass-balance–climate relationship. Due to the availability of bedrock maps of Mýrdalsjökull and Öraefajökull (Björnsson et al., 2000; Magnússon et al., 2012), further studies may include the development of fully coupled mass balance–ice dynamic models which could reproduce the series of dDEMs obtained for these glaciers in this study.

7. CONCLUSIONS

This study presents a 70-year record of elevation changes and geodetic mass balance of glaciers distributed across all parts of Iceland (excluding the three largest ice caps), most of them previously lacking mass-balance measurements. The glaciers were close to equilibrium during the period 1960–1994, with an area-weighted mass balance of 0.07 ± 0.07 m w.e. a^{-1} . This was followed by a rapid decline of the glaciers: -1.20 ± 0.09 m w.e. a^{-1} and a mass loss of 21.4 ± 1.6 Gt (1.3 ± 0.1 Gt a^{-1}), or 0.059 ± 0.005 mm SLE, during the period 1994–2010.

The region-wide, multitemporal intercomparison of mass balance revealed spatial patterns across Iceland: glaciers located

close to the south and west coast experience larger decadal oscillations in mass balance, a consequence of larger static mass-balance sensitivities to summer temperature and winter precipitation, than the interior, northern and eastern glaciers. This pattern can likely be explained by different local climate, related to oceanic currents surrounding Iceland, rain shadows, and elevation of the glaciers. Due to a large intraregional variability, particular care should be taken when extrapolating mass balance from one glacier to another, even at close distances.

A linear model relating mass balance, summer temperature and winter precipitation explained more than 70% of the observed mass-balance variability. Yet we acknowledge some limitations of this model, such as other mass balance forcing or the assumption of linearity between area and volume changes. The latter assumption was not satisfied at specific time periods, with lowering in the accumulation area while the glacier fronts were advancing. This was attributed to changes in ice flux toward the ablation area, suggesting additional complexity due to the dynamical response to mass-balance fluctuations. This encourages further studies aiming at coupling mass balance to ice dynamics, especially focused at reproducing events of increased ice flux toward the ablation area.

DATA AVAILABILITY STATEMENT

The mass-balance records calculated in this study are presented in **Supplementary Material**, and have been submitted to WGMS (www.wgms.ch). The glacier outlines will be uploaded to GLIMS [www.glims.org], Hannesdóttir et al., submitted. The DEMs and orthoimages created from aerial photographs, Hexagon KH9 photographs, SPOT 5 and ASTER, as well as the lidar DEMs, are available upon request. Maps of elevation difference using Pléiades data are available upon request.

AUTHOR CONTRIBUTIONS

JB, EM, and EB designed the research study and methods. JB performed the data processing and calculations. ÁG contributed in processing data for Mýrdalsjökull and Tungnafellsjökull. TJ prepared meteorological data for the statistical analysis of mass balance. FP and HB provided the glaciological mass balance of Vatnajökull and Langjökull. TT provided the glaciological mass balance of Hofsjökull. All coauthors interpreted the results and helped writing the manuscript.

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database of the outlines of Icelandic glaciers. Bolli Pálmason is acknowledged for providing downscaled precipitation data from regular forecast runs of the IMO for 2016–2019. Ken Moxham is acknowledged for the English-language editing of the manuscript. Michelle Koutnik is thanked for fruitful discussions about potential applications from the datasets. MZ, TS, and MH are acknowledged for their valuable comments during the revision of the manuscript. This study is based on the last chapter of the PhD dissertation of Belart (2018).

SUPPLEMENTARY MATERIAL

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/feart.2020.00163/full#supplementary-material>

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Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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